



# **Mathematical Modelling of Surface-Groundwater Interactions under Varying Hydrological Conditions: Review of Past Works**

**Jyoti Chetan Vanikar<sup>1\*</sup>, Rajeev Kumar Bansal<sup>2</sup> and Vineeta Basotia<sup>1</sup>**

<sup>1</sup>*JJT University, Jhunjhunu, Rajasthan, India.*

<sup>2</sup>*National Defence Academy, Khadakwasla, Pune, Maharashtra, India.*

## **Authors' contributions**

*This work was carried out in collaboration among all authors. Author JCV designed the study, performed the literature survey, wrote the protocol and wrote the first draft of the manuscript. Author RKB managed the analyses of the study. Author VB helped in preparation of the manuscript. All authors read and approved the final manuscript.*

## **Article Information**

DOI: 10.9734/JERR/2021/v20i417298

### Editor(s):

(1) Dr. Heba Abdallah Mohamed Abdallah, National Research Centre, Egypt.

### Reviewers:

(1) Siti Zulaiha Ahmad, Universiti Teknologi MARA (UiTM), Malaysia

(2) Oloro Obarhire John, Delta State University, Nigeria.

Complete Peer review History: <http://www.sdiarticle4.com/review-history/66562>

**Received 10 January 2021**

**Accepted 16 March 2021**

**Published 24 March 2021**

**Review Article**

## **ABSTRACT**

Surface-groundwater interaction is a research area of significant importance for its central role in wastewater treatment, irrigation, drainage, flood control, erosion and sediment control. Mathematical models are often used for the estimation of surface-groundwater interactions under the variety of hydrological conditions. Due to cost effectiveness and ability to accommodate variations in aquifer parameters, mathematical models have gained immense importance in the past few decades. The objective of this review paper is to portray the contribution of the hydrologist towards the growing area of surface-ground water interaction from all over the world who proposed, analyzed, executed and validated the developed Mathematical models. To begin with, we briefly introduce the main mathematical equations that govern the flow of groundwater in unconfined and confined aquifer systems. The development of stream-aquifer models is presented in a chronological order to provide a clear understanding of the contributions of past works. The methodology used in the past work is adequately discussed without going into mathematical

\*Corresponding author: Email: [jyotivanikar@gmail.com](mailto:jyotivanikar@gmail.com);

details. Furthermore, we also summarize recent developments concerning groundwater flow in presence of vertical streambed, partial penetration, stream-stage variations and multiple recharge/discharge basins.

*Keywords: Surface-groundwater; ditch-drain; sloping aquifer; water table; boussinesq equation.*

## 1. INTRODUCTION

Surface-groundwater models are gaining immense importance in the past few decades due to applications in the assessment of baseflow, conjunctive management of groundwater resources, catchment hydrology, recharging and dewatering of aquifer and solute transport in the coastal aquifer system. Estimation of surface-groundwater interaction is also important in artificial recharge scheme, site remediation and irrigation system. Mathematical models have emerged as efficient and cost-effective tools to obtain the quantitative approximation of the interaction between groundwater resources and hydrologically connected water bodies. Analytical models are also useful in understanding the interplay between various aquifer parameters and hydraulic properties of the aquifer [1].

## 2. BRIEF HISTORY OF HYDROLOGICAL DEVELOPMENT

The relation between hydrostatic and hydrodynamic theory was well defined by physicists, mathematicians and scientist by the eighteenth century. Danial Bernoulli (1700-1782) was the first among them to show that in steady, incompressible inviscid flow the energy is conserved along a streamline. This is represented as:

$$\frac{v^2}{2g} + h + z = \text{constant} \quad (1)$$

Where  $h$  =depthofthefluid,  $z$  =potentialhead,  $v^2/2g$ = kinetichead.

In 1823, a differential relationship for velocity and pressure in unsteady and three-dimensional viscous flow was represented by Claude-Louis Navier (1785-1836). With the additional work, it resulted into the famous Navier-Stokes equation which serves as the mathematical basis for ideal fluid flow. Navier-Stokes equation forms the backbone of fluid dynamics and hydrology. The year 1860 can be remarked as the origin of the

hydrological modelling. Later, Henry Darcy [2] performed an experiment of groundwater flow through sandy particles and established famous empirical formula known as Darcy's law which is a milestone in groundwater quantitative hydrology. Darcy's law states that the rate flow  $Q$  of water in a porous media is directly proportional to the cross sectional area  $A$  of the porous medium and piezometric head difference  $Q_1 - Q_2$  between two points and inversely proportional to the length  $L$  of process media, i.e.

$$Q = K \frac{A(Q_1 - Q_2)}{L} \quad (2)$$

Where  $K$  is the hydraulic conductivity of the porous media. The basic equation of groundwater flow can be derived by applying the principle of mass conservation on a representative elementary volume. The resulting equation can be expressed as

$$\frac{\partial}{\partial x} \left( K_x \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left( K_y \frac{\partial h}{\partial y} \right) + \frac{\partial}{\partial z} \left( K_z \frac{\partial h}{\partial z} \right) = S_s \frac{\partial h}{\partial t} \quad (3)$$

Where  $K_x, K_y, K_z$  are components of the hydraulic conductivity.  $h$  is the water head distribution and  $S_s$  is the specific yield.

Boussinesq [3] derived the nonlinear equation by combining Darcy's law with the mass balance equation which characterizes the governing equation of groundwater flow in most of the research pertaining to surface seepage flow. In order to accelerate the knowledge of groundwater flow mechanism, extensive innovations of mathematical modeling has been successively carried out for last one and half century. The earliest Mathematical model was developed in 1860. It was built on the basis of the Dupuit-Forchheimer hypothesis. In 1877 Bousinesq using method of separation of variable to obtain a general solution for governing equation of unconfined ground water flow on a bedrock in a plane with a steep hillslope. This mathematical approach was based on the assumption that the groundwater flow is almost parallel to the sloping bedrock. Theis [4] developed the relation between dropdown of

piezometric head and the discharge rate and period of discharge from a well. This paper presents an analytical solution of an unsteady groundwater flow in a confined aquifer. This solution is based on Theis [4] function analogy. Subsequent developments in this field under category includes with and without storage from the confined aquifers. The research of Hantush and Jacob [5] throws light on the above fact. They obtained an analytical solution of transient radial ground water flow in confined leaky aquifer system under withdrawal of water from aquifer storage. This process is advanced with the finite or an infinite extent of considered domain.

An analytical solution of one dimensional groundwater flow for confined and homogeneous aquifer is developed by Hantush [6] Cooper and Rorabaug [7], Hall and Moench [8], Tolikas et al. [9]. Polubarinova-Kochina [10] developed a series solution of the Boussinesq equation for transient ground water flow. The innovations by Ibrahim and Brutsaert [11] and Verma and Brutsaert [12] assesses the reliability and limitations of Dupuit-Forchheimer approach. Wooding [13] developed an analytical solution for saturated flow over slanting bed employing conformal mapping technique. By applying extended Dupuit-Forchheimer assumption, the approximation solution for ground water flow over inclined impervious base is established by Child [14]. Marino [15] presented an analytical solution for the prediction of water table variations induced by uniform and time dependent recharge in ditch drain. Using vertical and horizontal axes for defining the governing ground water flow, a simplified version of Child [14] approximation is developed by Chapman [16]. Beven [17] used kinematic wave approach for deriving a solution for saturated and unsaturated flows. Zecharies and Brutseart [18] disproved the kinematic wave approach for aquifers possessing very small slopes. By applying Laplace transform method Verhoest et al. [19] presented the analytic solution of a transient ground water flow over a sloping unconfined aquifer which is applicable for time varying recharge rate as it exists in natural environment. By considering nonlinear equation, the mechanism of saturated subsurface flow over a hillslope of uniform width is analysed by many researcher Chapman [20] and Basha and Maalouf [21].

The approximation solution for the linearised version of Boussinesq equation is developed by Sanford et al. [22], Brutseart [23], Su [24] etc.

Influence of varying recharge rate over water table fluctuations for distinct one dimensional and two dimensional flow is innovated by Singh and Rai [25], Rai and Singh [26-28], Rai et al. [29]. Effects of transient recharge over water table fluctuations in one dimensional sloping aquifer system is analysed by Ram and Chauhan [30], Singh et al. [31]. Innovations related to localised recharge over an inclined base is presented by Marino [32], Bansal [33]. An analytical solution of the water table fluctuations in two-dimensional aquifer system of finite extent is presented by Ramana et al. [34,35]. The prediction of water table variations induced by a single recharge basin at constant rate over different geometrical shapes is presented in Marino [36] Marino [15] and Teloglou [37].

Mathematical model explaining growth of the water table induced by exponentially decaying recharge rate over circular shape unconfined aquifer subjected to Dirichlet boundary condition is depicted in Rai et al. [38]. Some studies present the snapshot of three-dimensional model in unconfined aquifer of finite and infinite thickness [39,40]. The three dimensional numerical model built on the basis of Richards equation and ADE (Advection Dispersion equation) to simulate the effect storm water penetrating basin over its aquifer is presented by Bahar et al. [41]. Butler and Zlontik [42] developed a solution for estimating withdrawal rate and stream depletion rate occurred due to pumping. The expression for the water head and flow rate of an inclined aquifer adjacent to the water bodies is presented by Zlontik and Haung [43], Upadhyay and Chauhan [44-46]. Bansal and Das [47-49], Verhoest et al. [12] developed analytical models for analysing stream aquifer interaction over sloping aquifer subjected to distinct hydrological constraints. The estimation of water head, flow rate and steady state profile of a groundwater flow in a homogeneous unconfined sloping aquifer subjected to seepage from adjacent streams of varying water level and constant recharge is presented by Vanikar and Bansal [50]. Singh and Jaiswal [51] presented a numerical solution of two-dimensional free flow of water subjected to time varying recharge to aquifers underlain by a slanting impervious base. This work also depicts the picture of height dependent evapotranspiration impact on hydrological parameters. The unique model by considering Cauchy's boundary condition and Robin boundary condition is developed by Teloglou and Bansal [52] and Moutsopoulos [53]

subjected to different modes of water variations level. Numerical solution of the ground water flow by discretising spatial variable and temporal variable are obtained by using different techniques as finite difference method [54-57] and fourth order Runge-Kutta method [58,59]. Finite volume method is applied to obtain solution of two-dimensional unsaturated flow in regions of irregular shapes, for the prediction of two-dimensional transient ground water flow over unconfined slanting aquifer [60] to estimate depletion rates of a stream caused due to well pumping [61]. Seedpanah et al. [62] studies tidal fluctuations induced by ground water flow in coastal aquifers for studying ground water flow mechanism in porous medium of heterogeneous type. Three-dimensional analytical studies of the ground water flow is carried out by Haunget al. [63] and Chang et al. [64].

A digital numerical model was developed by Marino [65] applying predictor-corrector scheme for solving uni-dimensional Boussinesq equation in unconfined horizontal aquifer; however, Parlange [66] presented the solution by using partial differential equation solver PDE2D. The solution of one-dimensional Boussinesq equation of dimensionless format induced by constant recharge over unconfined sloping aquifer is derived by Beven [67]. Using finite difference method Renu et al. [68] developed numerical model for the dissolution of benzene and to study the transportation of aqueous state benzene in a fracture-matrix system of saturated nature subjected to steady-state ground water flow condition. The SUTRA [69], FEFLOW [70], MARUN [71] and SEWAT [72] are important benchmark of hydrology. Applying variational iteration method [73,74] developed analytical solution of nonlinear equation. A method of decomposition due to Adomian [75] proved to be efficient and accurate method for solving specially nonlinear ordinary and partial differential equations, integral equations, differential delay equations from engineering and applied sciences. The research work in this area is expanded by Wazwaz [76]. He developed new method, modified decomposition method. Duan [77] derived recurrence triangle formula for calculating polynomial. This work is followed by deduction of new algorithm for calculation of multivariable Adomian polynomials. The combination of these polynomials with other iterative techniques have been used for the approximation of nonlinear term of partial differential equations [78]. When these

polynomials are accomplished by differential transform method via algebraic recurrence relation results in Taylor series solution however application of integral operator inversely with method selected from homotopy analysis method (HAM) results in series solution.

Attili [79] depicted the solution of Solitons by Adomian decomposition method. It is of great importance in the field of fluid dynamics, magneto-hydro dynamics and also in water wave studies. The solution of one-dimensional wave equation called as acoustic equation is presented by Dispini et al. [80] by using Adomian decomposition method. The classical study of ground water flow developed by sudden change in piezometric head of semi-infinite aquifer is presented by Moutsopoulos [81] by using Adomian Decomposition method. The numerical analysis of nonlinear ordinary differential equation of second order by using Adomian Decomposition method is presented by Agom et al. [82] where continued algorithm is implemented in discrete domain. This principal is adopted in Maple package. Next remarkable development is seen in newborn approaches and tools for analysing hydrological observations and data. Even though the development of these approaches cannot be traced in hydrology but can be adapted for applications in hydrology. Some of these approaches involve artificial neural networks [83,84], time series analysis [85], geostatistical approach [86], time series by employing FORTRAN95 [87], data assimilation approach, validation and calibration techniques [88]. The mathematical model based on Finite Differenc Techniques for forecasting the variations in the groundwater level and pressure water level is proposed by Daliev [89]. These tools and techniques helps in better understanding of hydrologic system and hydrologic processes. The above mentioned mathematical models widely cover different dimensional i.e. one, two, three dimensional flows and radial flow, distinct types of aquifer domains as finite, infinite and semi-infinite domain. These models are developed for different boundary condition as Dirichlet boundary condition, Robin boundary condition as well as no boundary conditions. By employing mathematical techniques and numerical procedure, the solutions of these model are developed. In 1960s, with the introduction of computer and their amazing computational techniques and calculative power, the hydrological field gain a quantum leap in 1970s and 1980s and gave birth to digital hydrology or

numerical hydrology. The history of hydrological development is flooded with girth of extensive work in this field however this review depicts the picture of few important developments occurred in last two to two and half decades. For easy reference these milestones of hydrology have been arranged theme wise.

### 3. CONCLUSIONS

The hydrological development of the groundwater flow in relation with the aquifer systems for last two centuries from first quarter of 19th century to first quarter of 21th century are reviewed. Numerous earlier studies reveal that the studies of the groundwater flow play a major and active role in developing the solutions pertaining to the arising problems from ecology, engineering, hydrogeology, engineering, and environmental areas. The initial period can be regarded as conceptual period where development is based on concept and relationship between parameters. The middle half of the 19<sup>th</sup> century witnessed empirical approach which was necessity for establishing complex hydrological engineering projects as they were increasing number wise. This followed by flood estimating hydrological models. The birth of subsurface flow models was observed in the second half of 19 century. This period also witnessed the use of statistical methods to estimate the impact of storage on ground water flows as well as development of models using digital computer which is termed as digital hydrology or numerical hydrology. The later part of the above period shows the integration of hydrology with other branches such as ecohydrology, hydrogeology, coastal hydrology, hydro climatology, social hydrology etc. By considering the focus of above all aspects it's clear that groundwater is a multitasker.

### COMPETING INTERESTS

Authors have declared that no competing interests exist.

### REFERENCES

1. Bear J. Dynamics of fluids in porous media. Elsevier. New York; 1972.
2. Darcy H. Les fontaines publiques de la ville de Dijon. Paris: Dalmont; 1956.
3. Boussinesq J. Recherches théoriques sur l'écoulement des nappes d'eau infiltrées dans le sol et sur le débit de sources. J. Math. Pures Appl. 1904;10:5–78.
4. Theis CV. The relation between the lowering of the piezometric surface and the rate and duration of discharge of a well using ground-water storage. Trans Am Geophys Union. 1935;16:519–524.
5. Hantush MS, Jacob CE. Nonsteady radial flow in an infinite leaky aquifer. Trans Am Geophys Union. 1955;36:95–100.
6. Hantush MS. Drawdown around a partially penetrating well. J Hydraul Div Am Soc. 469 Civ. Eng. 1961;87:83–98.
7. Cooper-Jr HH, Rorabaugh MI. Groundwater movements and bank storage due to flood stages in surface streams. USGC water supply paper. 1963;1536-J:343-366.
8. Hall FR, Monech AF. Application of convolution equation to stream-aquifer relationship. Water Resource. Res. 1972; 8(2):487-493.
9. Tolikas PK, Sidiropoulos EG, Tzimopolous CD. A simple analytical solution for the Boussinesq one-dimensional groundwater flow equation. Water Resour. Res. 1984; 20(1):24–28.
10. Polubarinova-Kochina PY. Theory of Groundwater Movement (Trans: deWiest, J.M.R.) Princeton University Press, Princeton; 1962.
11. Ibrahim HA, Brutsaert W. Inflow hydrographs from large unconfined aquifers, Proc. Irrig Drain Div Am Soc Civ Eng. 1965;91:21–38.
12. Verhoest NEC, Pauwels VRN, Troch PA, Troch FPD. Analytical solution for transient water table heights and outflow from ditch-drained terrains. J Irrig Drain Eng. 2002; 128(6):358–36476.
13. Verma RD, Brutsaert W. Unsteady free surface ground water seepage. J Hydraul Div Am Soc Civ Eng. 1971;97(8):1213–1229.
14. Wooding RA, Chapman TG. Groundwater flow over a sloping impermeable layer: 1. Application of the Dupuit-Forchheimer assumption. J Geophys. Res. 1966;71: 2895–2902.
15. Childs EC. Drainage of groundwater resting on sloping beds. Water Resour Res. 1971;7:1256–1263.
16. Marino MA. Growth and decay of groundwater mounds induced by percolation. J Hydrol. 1974;22:295–301.

16. Chapman TG. Modeling groundwater flow over sloping beds. *Water Resour Res.* 1980;16(6):1114–1118.
17. Beven K. On subsurface stormflow: Predictions with simple kinematic theory for saturated and unsaturated flows. *Water Resources Research.* 1982;18(6).
18. Zecharias YB, Brutsaert W. Recession characteristics of groundwater outflow and base flow from mountainous watersheds; 1988.  
Available: <https://doi.org/10.1029/WR024i010p01651>
19. Verhoest NEC, Troch PA. Some analytical solutions of the linearized Boussinesq equation with recharge for a sloping aquifer. *Water Resour Res.* 2000;36(3): 793–800.
20. Chapman TG. Recharge-induced groundwater flow over a plane sloping bed: Solutions for steady and transient flow using physical and numerical models. *Water Resour. Res.* 2005;41.
21. Basha HA, Maalouf SF. Theoretical and conceptual models of subsurface hillslope flows. *Water Resource* 2004;41(7): W077018
22. Sanford WE. Hillslope drainage with sudden drawdown: Closed form solution and laboratory experiments. *Water Resources Research.* 1993;29(7):2313-2321.
23. Brutsaert W. The unit response of groundwater outflow from a hillslope. *WaterResour Res.* 1994;30(10):2759-2763
24. Su N. A formula for computation of time-varying recharge of ground-water. *J Hydrol.* 1994;160:123–135.
25. Singh RN, Rai SN. On subsurface drainage of transient recharge. *J Hydrol.* 1980;48:303–311.
26. Rai SN, Singh RN. A mathematical model of water table fluctuations in a semi-infinite aquifer induced by localized transient recharge. *Water Resour Res.* 1981;17(4): 1028–1032
27. Rai SN, Singh RN. Water table fluctuations in response to time varying recharge. *Scientific basis for water resources management, IAHS.* 1985;153:287–294.
28. Rai SN, Singh RN. Water table fluctuations in an aquifer system due to time varying surface infiltration and canal recharge. *J Hydrol.* 1992;136:381–387.
29. Rai SN, Manglik A, Singh RN. Water table fluctuation in response to transient recharge from a rectangular basin. *Water Resour Manag.* 1994;8(1):1–10.
30. Ram, Chauhan. Analytical and experimental solutions for drainage of sloping lands with time-varying recharge; *Water Resources Research.* 1987;23(6): 1090-1096.  
DOI: 10.1029/WR023i006p01090
31. Singh RN, Rai SN, Ramana DV. Water table fluctuation in a sloping aquifer with transient recharge. *J Hydrol.* 1991;126: 315–326.
32. Marino MA. Dynamic response of aquifer system to localized recharge 1 JAWRA. *Journal of the American Water Resources Association.* 1976;12(10):49-63.
33. Bansal RK. Unsteady seepage flow over sloping beds in response to multiple localized recharge. *Appl Water Sci;* 2015.  
DOI: 10.1007/s13201-015-0290-2
34. Ramana DV, Rai SN, Singh RN. Water table fluctuation due to transient recharge in a 2-D aquifer system with inclined base. *Water Resources Management.* 1995;9: 127–138.
35. Ramana DV, Rai SN, Singh RN. Water table fluctuation due to transient recharge in a 2-D aquifer system with inclined base. *Water Resour Manag.* 1995;9:127–138.
36. Marino MA. Hele-Shaw model study of the growth and decay of groundwater ridges *Journal of Geophysical Research Atmospheres.* 1967;72(4).  
DOI: 10.1029/JZ072i004p01195
37. Teloglou IS, Zissis TS, Pangopoulos AC. Water table fluctuation in aquifers overlying a semi-impervious layer due to transient recharge from a circular basin. *Journal of Hydrology.* 2008;348:215-223.
38. Rai SN, Singh RN. Evolution of the water table in a finite aquifer due to transient recharge from two parallel strip basins. *Water Resour Manag.* 1998;12:199–208
39. Dagen G. Linearized solution of free surface groundwater flow with uniform recharge, *J. Geophys. Res.* 1967;72:1183-1193.
40. Zlotnik V, Ledder G. Groundwater velocity in an unconfined aquifer with rectangular areal recharge. *Water Resources Research.* 1993;29(8):2827-2834.
41. Bahar T, Oxarango L, Castebrunet H, Rossier Y, Blondin FM. 3D modelling of solute transport and mixing during managed aquifer recharge with an

- infiltration basin. *Journal of Contaminant Hydrology*. 2021;237:103758.
42. Zlotnik VA, Zurbuchen BR, Ptak T. The steady-state dipole-flow test for characterization of hydraulic conductivity statistics in a highly permeable aquifer; Horkheimer Insel site, Germany, *Ground Water*. 2001;39(4):504–516.
  43. Zlotnik VA, Huang H. Effect of partial penetration and streambed sediments on aquifer response to stream stage fluctuations, *Ground Water*. 1999; 37(4):599–605.
  44. Upadhyaya A, Chauhan HS. An analytical solution for bi-level drainage design in the presence of evapotranspiration. *Agric. Water Manage.* 2000;45(2):169–184.
  45. Upadhyaya A, Chauhan HS. Interaction of stream and sloping aquifer receiving constant recharge. *J Irrig Drain Eng.* 2001a;127(5):295–301.
  46. Upadhyaya A, Chauhan HS. Falling water tables in non-sloping/sloping aquifer. *J Irrig Drain Eng.* 2001b;127(6):378–384.
  47. Bansal RK, Das SK. The effect of bed slope on waterhead and flow rate at the interfaces between the stream and groundwater: Analytical study. *J. Hydrol. Eng.* 2009;14(8):832–838.
  48. Bansal RK, Das SK. An analytical study of water table fluctuations in unconfined aquifers due to varying bed slopes and spatial location of the recharge basin. *J. Hydrol Eng.* 2010;15(11):909–917.
  49. Bansal RK, Das SK. Response of an unconfined sloping aquifer to constant recharge and seepage from the stream of varying water level. *Water Resour Manag.* 2011;25:893-911.
  50. Vanikar JC, Bansal RK. Mathematical modeling of subsurface seepage flow over sloping terrain due to vertical recharge and seepage from stream stage variations. *International Journal of Future Generation Communication and Networking*. 2020; 13(3s):1241–1248
  51. Singh S, Jaiswal CS. Numerical Solution of 2D Free Surface to Ditch Drains in Presence of Transient Recharge and Depth-Dependent ET in Sloping Aquifer; *Water Resource Management*. 2006;20(5): 779-793.
  52. Teloglou IS, Bansal RK. Transient solution for stream–unconfined aquifer interaction due to time varying stream head and in the presence of leakage. *J. Hydrol.* 2012;428–429:68–79.
  53. Moutsopoulos KN. Solution of boussinesq equation subject to a nonlinear Robin boundary condition. *Water Resources Research*. 2013;49:7-18. DOI: 10.1029/2012WR012221,2013
  54. Lin CL. Digital simulation of the Boussinesq equation for a water table aquifer. *Water Resour. Res.* 1972;8(3): 691–698.
  55. Bear J, Verruijt A. *Modeling groundwater flow and pollution*. D. Reidel Publishing Company. 1987;414.
  56. Walton WC. *Numerical Groundwater Modeling*. Lewis Publishers, Chelsea, MI. Waterloo Hydrogeologic, 2004. Visual MODFLOW. Waterloo Hydrogeologic Inc, Waterloo, Ontario, Canada; 1989.
  57. Anderson MP, Woessner WW. *Applied groundwater modeling*. Academic Press, Inc., San Diego, CA. 1992;381.
  58. Koussis A. An analytical-numerical solution of transient flow through nonsaturated porous media; conference: 16th midwestern mechanics conference at: Kansas State University, Manhattan, Kansas. 1979;183-187.
  59. Yeh WWG. Nonsteady flow to surface reservoir. *J Hydraul. Div. ASCE HY.* 1970;3:609-618.
  60. Kalaidzidou PN, Karamouzis D, Moraitis D. A finite element model for the unsteady groundwater flow over sloping beds. *Water Resour Manag.* 1997;11(1):69–81.
  61. Darama. An analytical solution for stream depletion by cyclic pumping of wells near streams with semipervious beds; 2001. Available: <https://doi.org/10.1111/j.1745-6584.2001.tb00353.x>
  62. Saeedpanah I, Jabbari E, Shayanfar MA. Numerical simulation of groundwater flow via a new approach to the local radial point interpolation meshless method. *Int J Comput Fluid D.* 2011;25(1):17–30.
  63. Huang CS, Chen JJ, Yeh HD. Approximate analysis of three-dimensional groundwater flow toward a radial collector well in a finite –extent unconfined aquifer. *Hydrol. Earth Syst Sci.* 2016;20:55-71. DOI: 10.5194/hess-20-55-2016.
  64. Chang CH, Huang CS, Yeh HD. Analysis of three dimensional unsaturated-saturated flows induced by localized recharge in unconfined aquifers. *Hydrol Earth Syst Sci.* 2017;22:3951–3963.

65. Marino MA. Digital simulation model of aquifer response to stream stage fluctuation. *Journal of Hydrology*. 1975; 25(1):51-58.
66. Parlange JY, Hogarth WL, Govindaraju RS, Parlange MB, Lockington D. On an exact analytical solution of the Boussinesq equation, *Transport Porous Media*. 2000; 39:339–345.
67. Beven K. Kinematic subsurface stormflow. *Water Resour Res*. 1981;17(5):1419–1424.
68. Renu V, Suresh Kumar G. Numerical modeling on benzene dissolution into groundwater and transport of dissolved benzene in a saturated fracture-matrix system environmental processes volume. 2016;3:781–802.
69. Voss CI. Sutra – A finite element simulation model for saturated – unsaturated fluid density dependent groundwater flow with energy transport or chemically single species contaminant transport. USGS Water Resources, Investigate Report. 1984;84-4269:409.
70. Diersch HG. Feflow; interactive, graphics-based finite-element simulation system for modeling groundwater flow, contaminant mass and heat transport processes. WASY Institute for Water Resources Planning and System Research Ltd; 1998.
71. Boufadel MC, Suidan MT, Venosa AD, Bowers MT. Contribution of capillary flow to steady seepage: application to trenches and dams. *J Hydraulic Engng ASCE*. 1999b;125:286-294.
72. Langevin CD. SEAWAT: A computer program for simulation of variable-density groundwater flow and multi-species solute and heat transport: U.S. Geological Survey Fact Sheet. 2009;3047:2.
73. He JH. Variational iteration method - A kind of non-linear analytical technique: Some examples. *International Journal of Non-Linear Mechanics*. 1999;34(4):699-708.  
DOI: 10.1016/S0020-7462(98)00048-1
74. Wu, He. Construction of solitary solution and compacton-like solution by variational iteration method *Chaos Solitons & Fractals*. 2006;29(1):108-113.  
DOI: 10.1016/j.chaos.2005.10.100
75. Adomian G. Solving Frontier problems of Physics. *The Decomposition Method. of Fundamental Theories of Physics*, Kluwer Academic Publishers, Boston, mass, HUSA. 1994;6.
76. Wazwaz MA. A reliable modification of adomian decomposition method. *Appl Math Comput*. 1999;102:77-86.
77. Duan JS. Recurrence triangle for Adomian polynomials, *Appl Math Comput*. 2010; 216:1235–1241.
78. El- Wakil SA, Abdou MA. Modified extended tanh-function method for solving nonlinear partial differential equations physics letters A. 2007;299(2):179-188.  
DOI: 10.1016/S0375-9601(02)00669-2
79. Attili BS. The adomian decomposition method for solving boussinesq equation arising in water wave propagation. Wiley Inter Science; 2006.  
DOI: 10.1002/num.
80. Dispiniot M, Mungasi S. Adomian decomposition method used to solve the one-dimensional acoustic equations. *Journal of Physics: Conference Series*. 2017;856(1).  
Article id: 012003.
81. Moutsopoulos KN. The analytical solution of the Boussinesq equation for flow induced by a step change of the water table elevation revisited, *Transp. Porous Media*. 2010;85(3):919–940.
82. Agom E. Application of adomian decomposition method in solving second order nonlinear ordinary differential equations. *International Journal of Engineering Science Invention*. 2015;4(11): 60-65.
83. Niar SS, Sindhu G, et al. 2016. Groundwater level forecasting using Artificial Neural Network. *International Journal of Scientific and Research Publications*. 2016;6(1).
84. Tayfur G, Singh VP. Artificial neural networks Chapter 11. In: Singh VP(ed) *Handbook of applied hydrology*. McGraw-Hill Education, New York. 2017;11-1–11-6.
85. Bakker M, Schaars F. Solving groundwater flow problems with time series analysis: You may not even need another model. 2019;57(6):826–833.  
DOI: 10.1111/gwat.12927
86. Dwivedi D, Dafflon B, Arora B, Wainwright HM, Finsterle S. Spatial analysis and geostatistical methods. Chapter 20. In: Singh VP (ed) *Handbook of applied hydrology*. McGraw-Hill Education, New York; 2017.
87. Park E, Parker JC. A simple model for water table fluctuations in response to

- precipitation. Journal of Hydrology. 2008; 356:344-349.
88. Todini E, Biondi D. Calibration, parameter estimation, uncertainty, dataassimilation, sensitivity analysis and validation Chapter 22. In: Singh VP (Ed) Handbook of applied hydrology. McGraw-Hill Education, New York. 2017;22-1–22-19.
89. Daliev S. Mathematical modeling to change the groundwater level in the multilayer porous media. International Journal of Advanced Science and Technology. 2020;29(7):3366-3381.

---

© 2021 Vanikar et al.; This is an Open Access article distributed under the terms of the Creative Commons Attribution License (<http://creativecommons.org/licenses/by/4.0>), which permits unrestricted use, distribution, and reproduction in any medium, provided the original work is properly cited.

*Peer-review history:*  
*The peer review history for this paper can be accessed here:*  
<http://www.sdiarticle4.com/review-history/66562>